# A revamped optimally interpolated sea ice motion dataset

Charles Brunette<sup>1</sup>, Bruno Tremblay<sup>1</sup>, Robert Newton<sup>2</sup>

<sup>1</sup>McGill University, Montreal, Canada; <sup>2</sup>Lamont-Doherty Earth Observatory (Columbia), Palisades, USA

Corresponding author: <a href="mailto:charles.brunette@mail.mcgill.ca">charles.brunette@mail.mcgill.ca</a>







- From passive-microwave & radiometer imagery
  - SMMR, 1979-1987, 75 km
  - AVHRR, 1981-2000, 50 km
  - AMSRE, 2002-2011, 37.5 km
  - SSMI, 1987-2006, 75 km
  - SSMI/S, 2007-present, 75 km
- Good spatial coverage in the winter; less coverage in the summer months due to melt ponds, open water, and atmospheric water vapor

# **Optimal interpolation**

- Objective: combine sea ice motion vectors from different sources into a seamless dataset
- We retrieve the **buoy** and **satellite**-derived ice motion estimates from the *Polar Pathfinder* dataset (NSIDC, Tschudi et al. 2020), and introduce new **free drift** estimates (Brunette et al. 2021).
- Daily, 25 km resolution, merged ice motion vectors calculated as a weighted average, where the weights are based on:

#### Distance

- Inverse distance weighting options:  $e^{-d/r_o}$ ;  $1/x^p$ ;  $1 d/r_o$
- Maps of spatial correlation of  $u_i$  will define the local  $r_o$

#### > Error

<sup>2019.001</sup>

- Taking into account the source-based error (rmse)
- Including the spatial variation of the error (see map) and temporally evolving error (seasonality)

# Buoy data

- International Arctic Buoy Programme
- 1979-present, daily
- We take buoy drift as truth



## Free drift estimates

• First order momentum balance, sea ice motion in response to the atmospheric and oceanic stresses:  $\vec{u}_i = \alpha(h_i)\vec{u}_a + \vec{u}_o$ 

Atmospheric forcing from

E.g.: Spatial distribution of the error for the free drift estimates



 Resulting sea ice motion vectors will be validated against independent buoy data (ITPs).

### Outcomes

- Improved accuracy of sea ice velocity estimates
- Important innovation is expected in the summer months; with less satellitederived vectors, the merged ice velocity estimates rely more heavily on free drift.
- Support applications based on Lagrangian tracking, which include retrieving sea ice



- ERA5 10 m winds
- Climatology of the surface ocean current calculated as a residual
- Wind-ice transfer coefficient is linearly dependent on ice thickness (PIOMAS)
- 1979-present, daily, 25 km
- Ref: Brunette et al. (2021)



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Configuration:

age, investigating mechanisms for seasonal forecasting, informing socioenvironmental studies by quantifying pollutant transport, etc.

"Sometimes when there are strong winds, the new ice and the land-fast ice cannot come in contact with each other because the northerly winds cause the newly formed ice to break up and drift away. After the winds die down and the weather improves, the resultant open water freezes again and the current will move the new ice back and forth against the land-fast ice. [...] This is true, and that is the nature of the moving ice ." -Elder Aipilik Inuksuk, Igloolik, Nunavut

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